Supporting Information "A" for

Long-term phosphorus mass-balance of Lake Erie (Canada-USA) reveals a major contribution of in-lake phosphorus loading

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ID #	Study	Lake St. Clair (LSC)		Lake Erie (LE)			LES	Data Sources	Method
	Period		Western Basin	Central Basin	Eastern Basin	Whole Lake Erie	- (LSC +LE)		
1	P1	-	378 to 808	-	-	-	-	S 1	M1, M2
2	P2	-	-	300 to 1250	-	-	-	S2	M2
3	P3	-	-	2400 range: 0 to 4229 mean: 1872	-	-	-	S 3	M3
4	P4	-	415*	-	-	-	-	S4	M1
5	P5	-	-	10599	359	-	-	S5	M4
6	P6	-	-	10000-11000	-	-	-	S 6	M5
7	P7	220	495	795	2273	3563	3783	S 7	M6 (this study)
8	P7	219	479	763	2288	3530	3749	S 7	M7 (this study)
9	P7	220	480	706	2268	3454	3674	S 7	M8 (this study)

Table A1. Literature derived estimates of mean annual internal phosphorus load in metric tons (MT) in lakes St. Clair and Erie scaled up to the basin or whole lake areas. ID #1-5 indicate annual fluxes of gross internal load, while ID #6-8 indicate annual net internal load.

* Just for a four-day episode only (neither annul nor seasonal load).

- Study period: P1, 2013: Jun 24_& Jul 6, 2014: Jun 27 & Aug 4; P2, spring and summer of 2010 and 2012; P3, August-September of 1985 to 2012; P4, four-day episode in summer 2014; P5, long-term: analysis of *in-situ* collected sediment cores (30 cm long); P6, continuous in-situ nutrient analyzer measurements at 2 sites in summer 2019 and then scaling up to the basin size; P7, Oct 1, 2003 to Sep 30, 2016.
- <u>Data sources</u>: S1, Matisoff et al. (2016); S2, Paytan et al. (2017); S3, Nürnberg et al. (2019); S4, Gibbons & Bridgeman (2020); S5, Wang et al. (2021); S6, Anderson et al. (2021); S7, this study.
- <u>Methods</u>: M1, laboratory incubations scaled up to the basin size; M2, laboratory incubations of sediment cores scaled up to the basin size; M3, as a product of the estimated annual hypoxic factor (HF) values and a phosphorus release rate; M4, *in-situ* core profile analysis; M5, continuous *in-situ* nutrient analyzer measurements at 2 sites and then scaling up to the basin size; M6, <u>Net Mass Balance Approach</u> (NMBA) using steady state formulations (no change in basin concentrations between 2003 and 2016); M7, NMBA using non-steady state formulations (change in basin concentration $\neq 0$ and based on averages between April and August concentrations with difference determined between 2003 and 2016); M8, NMBA using non-steady state formulations (change in basin concentration between two 3-year periods 2003-2005 and 2014-2016).

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#		Quaternary Watersheds				Tertiary	Watersheds	
	ID #	Name	Drainage Area (km ²)	Receiving Waterbody*	ID #	Name	Latitude Longitude	Area (km ²)
1.	02GG-01	Upper St Clair River	186	SCR	02GG	Sydenham	42.7658	3496
2.	02GG-02	Lower St Clair River	316	SCR			82.0565	
3.	02GG-03	Lower Lake St. Clair Tributaries	88	LSC				
4.	02GG-04	Upper Lake St. Clair Tributaries	230	LSC				
5.	02GG-05	East Sydenham River	1505	LSC				
6.	02GG-06	Lower N. Sydenham River	550	LSC				
7.	02GG-07	Bear Creek	621	LSC				
8.	02GD-01	S. Thames River	764	LSC	02GD	Upper Thames	43.2121	3057
9.	02GD-02	Waubuno Creek	135	LSC			81.0424	
10.	02GD-03	Middle Thames River	294	LSC				
11.	02GD-04	Reynolds Creek	166	LSC				
12.	02GD-05	N. Thames River	734	LSC				
13.	02GD-06	Medway River	202	LSC				
14.	02GD-07	Fish - Flat Creeks	154	LSC				
15.	02GD-08	Trout Creek	173	LSC				
16.	02GD-09	Avon River	163	LSC				
17.	02GD-10	Black Creek	142	LSC				
18.	02GD-11	Whirl Creek	130	LSC				
19.	02GE-01	Lower Thames River	1607	LSC	02GE	Lower Thames	42.5548	2818
20.	02GE-02	Big Creek	251	LSC			81.8790	
21.	02GE-03	Jeanette Creek	416	LSC				
22.	02GE-04	McGregor Creek	285	LSC				
23.	02GE-05	Oxbow Creek	89	LSC				
24.	02GE-06	Dingham Creek	170	LSC				

*SCR, St. Clair River; LSC, Lake St. Clair; DR, Detroit River; WB, western basin; CB, central basin; EB, eastern basin.

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Table A2. (Continued)

#		Quaternary Watersheds			Tertiary Watersheds			
	ID #	Name	Drainage Area (km ²)	Receiving Waterbody*	ID #	Name	Latitude Longitude	Area (km ²)
25.	02GH-01	Pike Crk - Puce River	270	LSC	02GH	Cedar	42.1565	1765
26.	02GH-02	Ruscom River	210	LSC			82.7822	
27.	02GH-03	Belle River	150	LSC				
28.	02GH-04	Windsor Area - Little River	116	DR				
29.	02GH-05	River Canard	347	DR				
30.	02GH-06	Lower Detroit River	113	WB				
31.	02GH-07	Cedar Creek - Oxley - Seacliffe	228	WB				
32.	02GH-08	Sturgeon Point - Point Pelee	79	WB				
33.	02GH-09	Hillman - Lebo Creeks	206	CB				
34.	02GH-10	Pelee Island	41	WB				
35.	02GH-11	Fighting Island	5	DR				
36.	02GF-01	Renwick - Erie Beach	46	СВ	02GF	Rondeau	42.4741	770
37.	02GF-02	Flat Creek - Rondeau	115	CB			81.7300	
38.	02GF-03	Morpeth - Palmyra Beach	99	CB				
39.	02GF-04	Brock's Creek	213	CB				
40.	02GF-05	Tyconnel Beach	101	CB				
41.	02GF-06	Talbot Creek	196	CB				
42.	02GC-01	Stoney Creek	123	EB	02GC	Big Creek	42.8223	3992
43.	02GC-02	Kettle Creek	463	CB			80.6028	
44.	02GC-03	Catfish Creek	417	CB				
45.	02GC-04	Big Otter Creek	817	CB				
46.	02GC-05	Long Point	220	EB				
47.	02GC-06	South Otter - Clear Creeks	37	CB				
48.	02GC-07	Diedrich - Young Crks	249	EB				
49.	02GC-08	Big Creek	741	EB				

*SCR, St. Clair River; LSC, Lake St. Clair; DR, Detroit River; WB, western basin; CB, central basin; EB, eastern basin.

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Table A2. (Continued)

#		Quaternary Watersheds			Tertiary Watersheds			
	ID #	Name	Drainage Area (km ²)	Receiving Waterbody*	ID #	Name	Latitude Longitude	Area (km ²)
50.	02GC-09	Lynn River	301	EB				
51.	02GC-11	Nanticoke Creek	216	EB				
52.	02GC-12	Sandusk Creek	173	EB				
53.	02GC-13	Welland River	143	EB				
54.	02GA-01	Upper Grand River	1903	EB				
55.	02GA-02	Nith River	1131	EB			43.6352 80.4700	
56.	02GA-03	Galt Creek	108	EB	0264	Upper Grand		1715
57.	02GA-04	Speed River	505	EB	020A	Opper Oralid		4745
58.	02GA-05	Eramosa River	274	EB				
59.	02GA-06	Conestogo River	824	EB				
60.	02GB-01	Lower Grad River	684	EB			43.1122	
61.	02GB-02	McKenzie River	371	EB				
62.	02GB-03	Big Creek	166	EB	02GB	Lower Grand		2030
63.	02GB-04	Fairchield Creek	412	EB			00.1700	
64.	02GB-05	Horner Creek	397	EB				
65.	02HA-05	Upper Welland Canal	99	EB			42.0591	
65.	02HA-07	Welland River	1118	NR	02HA	Niagara River	79.4324	1407
67.	02HA-08	Fort Erie Creeks	190	NR, EB				
	TOTA	AL for St. Clair River	502					
	TOTA	AL for Lake St. Clair	9499					
	TOTAL for Detroit River		468					
	TOTAL for Lake Erie West Basin		431					
	TOTA	AL for Lake Erie Central Basin	2710					
	TOTA	AL for Lake Erie East Basin	9033					
	GRA	ND TOTAL	1407					

*SCR, St. Clair River; LSC, Lake St. Clair; DR, Detroit River; WB, western basin; CB, central basin; EB, eastern basin.

Table A3. Characterization of Lake Erie system watersheds located in United States and based on the 6-digit and 8-digit hydrological unit codes (HUC6 and HUC8).

#		Hydrologic Unit Coc	le 8-digit (HUC8)		Hydrologic Unit Code 6-digit (HUC6)			
	Code	Name	State	Watershed Area (km ²)	Code	Name	Watershed Area (km ²)	
1.	04090001	St. Clair	MI	2997.0	040900	St Clair - Detroit		
2.	04090002	Lake St. Clair	MI	662.8				
3.	04090003	Clinton	MI	2064.4				
4.	04090004	Detroit	MI	1518.3				
5.	04090005	Huron	MI	2378.3				
6.	04100001	Ottawa-Stony	MI	1806.3	041100	Western Lake Erie		
7.	04100002	Raisin	MI	2753.3				
8.	04100003	St. Joseph	OH	2832.3				
9.	04100004	St. Marys	OH	2054.8				
10.	04100005	Upper Maumee	OH	1003.3				
11.	04100006	Tiffin	OH	2014.0				
12.	04100007	Auglaize	OH	4316.1				
13.	04100008	Blanchard	OH	1999.4				
14.	04100009	Lower Maumee	OH	2798.6				
15.	04100010	Cedar-Portage	OH	2510.3				
16.	04100011	Sandusky	OH	4732.6				
17.	04100012	Huron-Vermilion	OH	1979.8				

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Table A3. (Continued)

#		Hydrologic Unit Code 8-digit (HUC8)				Hydrologic Unit Code 6-digit (HUC6)		
	Code	Name	State	Watershed Area (km ²)	Code	Name	Watershed Area (km ²)	
18.	04110001	Black-Rocky	OH	2324.7	041000	Southern Lake Erie		
19.	04110002	Cuyahoga	OH	2101.5				
20.	04110003	Ashtabula-Chagrin	OH	1641.5				
21.	04110003	Ashtabula-Chagrin	PA	156.6				
22.	04110004	Grand	OH	1827.5				
23.	04120101	Chautauqua-Conneaut	ОН	350.9	041201	Eastern Lake Erie		
24.	04120101	Chautauqua-Conneaut	PA	1571.5				
25.	04120101	Chautauqua-Conneaut	NY	868.0				
26.	04120102	Cattaraugus	NY	1449.5				
27.	04120103	Buffalo-Eighteenmile	NY	1856.1				
28.	04120104	Niagara	NY	2069.9				

#	WSC*	Gauge Name	Geographi	c Coordinates	Area	Tertiary
	Gauge		Latitude	Longitude	(km ²)	Watershed Codes
1	02GG002	Sydenham River Near Alvinston	42,8306	81 8517	701	0266
2.	02GG002	Sydenham River At Florence	42.6506	82.0083	1150	0200
3.	02GG006	Bear Creek Near Petrolia	42.9058	82.1189	249	
4.	02GG009	Bear Creek Below Brigden	42.8119	82.2983	536	
5.	02GG010	St. Clair River At Point Edward	42.9803	82.4022	n/a	
6.	02GG011	St. Clair River At Port Lambton	42.6572	82.5069	n/a	
7.	02GG013	Black Creek Near Bradshaw	42.7622	82.2592	213	
8.	02GE002	Thames River At Byron	42.9625	81.3317	3080	02GE
9.	02GE003	Thames River At Thamesville	42.5447	81.9672	4370	
10.	02GE007	McGregor Creek Near Chatham	42.3833	82.0950	204	
11.	02GH002	Ruscom River Near Ruscom Station	42.2114	82.6289	95	02GH
12.	02GH003	Canard River Near Lukerville	42.1589	83.0189	159	
13.	02GH005	Lake St. Clair At Belle River	42.2961	82.7108	n/a	
14.	02GH011	Little River At Windsor	42.3097	82.9283	55.3	
15.	02GC002	KETTLE CREEK AT ST. THOMAS	42.7775	81.2139	331	02GC
16.	02GC007	BIG CREEK NEAR WALSINGHAM	42.6856	80.5383	567	
17.	02GC008	LYNN RIVER AT SIMCOE	42.8233	80.2894	144	
18.	02GC010	BIG OTTER CREEK AT TILLSONBURG	42.8572	80.7233	354	
19.	02GC014	YOUNG CREEK NEAR VITTORIA	42.7656	80.2944	65.8	
20.	02GC018	CATFISH CREEK NEAR SPARTA	42.7458	81.0569	295	
21.	02GC021	VENISON CREEK NEAR WALSINGHAM	42.6533	80.5483	68.4	
22.	02GC022	NANTICOKE CREEK AT NANTICOKE	42.8097	80.0761	177	

Table A4. Characterization of the water gauging stations in Canadian part of the Lake Erie system.

*The Water Survey of Canada (WSC)

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Table A4. (Continued)

#	WSC*	Gauge Name	Geographi	c Coordinates	Area	Tertiary
	Gauge		Latitude	Longitude	(km ²)	Watershed Codes
	Number					
22.	02GC026	BIG OTTER CREEK NEAR CALTON	42.7106	80.8406	665	02GC
23.	02GC029	KETTLE CREEK ABOVE ST. THOMAS	42.8350	81.1347	134	
24.	02GC030	CATFISH CREEK AT AYLMER	42.7736	80.9825	127	
25.	02GC031	DODD CREEK BELOW PAYNES MILLS	42.7872	81.2675	99.6	
26.	02GC036	SILVER CREEK NEAR GROVESEND	42.6758	80.9533	40.3	
27.	02GB010	MCKENZIE CREEK NEAR CALEDONIA	43.0339	79.9497	173	02GB
28.	02GA003	GRAND RIVER AT GALT	43.3531	80.3156	3520	02GA
29.	14827	Grand River at York	43.0217	79.8913	6044	
30.	02HA019	WELLAND CANAL DIVERSION FROM LAKE ERIE	42.9500	79.2167		02HA
31.	02HA003	NIAGARA RIVER AT QUEENSTON	43.1569	79.0472	686000	
32.	02HA007	Welland River Below Caistor Corners	43.0217	79.6178	223	
33.	02HA013	NIAGARA RIVER AT FORT ERIE	42.9303	78.9142	683000	
34.	02HA024	Oswego Creek At Canboro	42.9911	79.6781	83.2	

*The Water Survey of Canada (WSC)

#	USGS	Gauge Name (State)	Geographic	Coordinates	Area	Hydrologic Unit
	Gauge		Latitude	Longitude	(km ²)	
	Number					
1.	04159130	St. Clair River At Port Huron (MI)	42°59'13"	82°25'29"	576014*	4090001
2.	04159492	Black River Near Jeddo (MI)	43°09'03"	82°37'29"	1202	4090001
3.	04159900	Mill Creek Near Avoca (MI)	43°03'16"	82°44'05"	438	4090001
4.	04160600	Belle River At Memphis (MI)	42°54'03"	82°46'09"	391	4090001
5.	04165500	Clinton River At Moravian Drive At Mt. Clemens (MI)	42°35'45"	82°54'32"	1901	4090003
6.	04165710	Detroit River At Fort Wayne At Detroit (MI)	42°17'53"	83°05'34"	592590*	4090004
7.	04166500	River Rouge At Detroit (MI)	42°22'23"	83°15'17"	484	4090004
8.	04167000	Middle River Rouge Near Garden City (MI)	42°20'53"	83°18'42"	259	4090004
9.	04168400	Lower River Rouge At Dearborn (MI)	42°18'30"	83°15'10"	236	4090004
10.	04168580	Ecorse River At Dearborn Heights (MI)	42°16'10"	83°17'23"	26	4090004
11.	04174500	Huron River At Ann Arbor (MI)	42°17'13"	83°44'02"	1888	4090005
12.	04174518	Malletts Creek At Ann Arbor (MI)	42°15'53"	83°41'18"	28	4090005
13.	04176500	River Raisin Near Monroe (MI)	41°57'38"	83°31'52"	2699	4100002
14.	04177000	Ottawa River At University of Toledo (OH)	41°39'35"	83°36'45"	388	4100001
15.	04193500	Maumee River At Waterville (OH)	41°30'00"	83°42'46"	16395	4100009
16.	04193999	Wolf Creek At Holland (OH)	41°36'34"	83°41'03"	64	4100009
17.	04195820	Portage River Near Elmore (OH)	41°29'28"	83°13'29"	1279	4100010
18.	04198000	Sandusky River Near Fremont (OH)	41°18'28"	83°09'32	3240	4100011

Table A5. Characterization of water gauging stations in US part of the Lake Erie system.

* including the upper Laurentian Great Lakes.

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Table A5. (Continued)

#	USGS	Gauge Name (State)	Geographic	Coordinates	Area	Hydrologic Unit
	Gauge Number		Latitude	Longitude	(km ²)	
19.	04199000	Huron River At Milan (OH)	41°18'03"	82°36'30"	961	4100012
20.	04199155	Old Woman Creek At Berlin Rd Near Huron (OH)	41°20'54"	82°30'50"	57	4100012
21.	04199500	Vermilion River Near Vermilion (OH)	41°22'55"	82°19'01"	679	4100012
22.	04200500	Black River At Elyria (OH)	41°22'49"	82°06'17"	1026	4110001
23.	04201500	Rocky River Near Berea (OH)	41°24'27"	81°52'58"	692	4110001
24.	04201526	Abram Creek At Kolthoff Drive At Brook Park (OH)	41°23'35"	81°51'01"	21	4110002
25.	04208504	Cuyahoga River Near Newburgh Heights (OH)	41°27'45"	81°40'52"	2041	4110002
26.	04208700	Euclid Creek At Cleveland (OH)	41°34'56"	81°33'32"	60	4110003
27.	04209000	Chagrin River At Willoughby (OH)	41°37'51"	81°24'13"	637	4110003
28.	04212100	Grand River Near Painesville (OH)	41°43'08"	81°13'41"	1774	4110004
29.	04213000	Conneaut Creek At Conneaut (OH)	41°55'37"	80°36'15"	453	4120101
30.	04213075	Brandy Run Near Girard (PA)	41°59'31"	80°17'29"	12	4120101
31.	04213152	Walnut Creek Upstream Pool Near Erie (PA)	42°04'26"	80°14'05"	97	4120101
32.	04213500	Cattaraugus Creek At Gowanda (NY)	42°27'48"	78°56'03"	1137	4120102
33.	04214500	Buffalo Creek At Gardenville (NY)	42°51'17"	78°45'18"	368	4120103
34.	04215000	Cayuga Creek Near Lancaster (NY)	42°53'24"	78°38'42"	250	4120103
35.	04215500	Cazenovia Creek At Ebenezer (NY)	42°49'47"	78°46'30"	350	4120103
36.	04216000	Niagara River At Buffalo (NY)	42°52'40"	78°54'59"	682980*	4120200
37.	04216220	Niagara River At Black Rock Lock At Buffalo (NY)	42°56'02"	78°54'17"	682980*	4120104
38.	04218000	Tonawanda Creek At Rapids (NY)	43°05'35"	78°38'10"	904	4120104
39.	04218518	Ellicott Creek Below Williamsville (NY)	42°58'40"	78°45'49"	211	4120104

* including the upper Laurentian Great Lakes.

#	Gauge #	Name (Location)	Latitude	Longitude	Other
					known
					ID #
I.	St. Clair River:				
1	9014098	Fort Gratiot (US, MI)	43° 0.40 N	82° 25.3 W	
2	9014096	Dunn Paper (US, MI)	43° 0.20 N	82° 25.3 W	
3	11940	Point Edward (CA, ON)	42° 59.5" N	82° 25.3 W	02GG010
4	9014090	Mouth Of The Black River (US, MI)	42° 58.5 N	82° 25.1 W	
5	9014087	Dry Dock (US, MI)	42° 56.7 N	82° 26.6 W	
6	9014080	St Clair State Police (US, MI)	42° 48.7 N	82° 29.1 W	
7	11950	Port Lambton (CA, ON)	42° 39.4 N	82° 30.4 W	02GG011
8	9014070	Algonac (US, MI)	42° 37.3 N	82° 31.6 W	
II.	Lake St. Clair:				
9	11965	Belle River (CA, ON)	42° 17.8 N	82° 42.7 W	02GH005
10	9034052	St Clair Shores (US, MI)	42° 28.4 N	82° 52.8 W	
III.	Detroit River:				
11	9044049	Windmill Point (US, MI)	42° 21.4 N	82° 55.8 W	
12	9044035	Fort Wayne (US, MI)	42° 17.9 N	83° 05.6 W	
13	9044030	Wyandotte (US, MI)	42° 12.1 N	83° 08.8 W	
14	11995	Amherstburg (CA, ON)	42° 08.7 N	83° 06.8 W	02GH008
15	9044020	Gibraltar (US, MI)	42° 05.5 N	83° 11.2 W	
16	12005	Bar Point (US, MI)	42° 03.7 N	83° 06.9 W	02GH009
IV.	Lake Erie:				
17	12065	Kingsville (CA, ON)	42° 01.6 N	82° 44.1 W	02GH010
18	12250	Erieau (CA, ON)	42° 15.6 N	81° 54.9 W	02GF002
19	12400	Port Stanley (CA, ON)	42° 39.5 N	81° 12.8 W	02GC027
20	12710	Port Dover (CA, ON)	42° 46.9 N	80° 12.1 W	02GC028
21	12865	Port Colborne (CA, ON)	42° 52.5 N	79° 15.2 W	02HA017
22	9044036	Fort Wayne (US, MI)	42° 17.9 N	83° 05.6 W	
23	9063090	Fermi Power Plant (US, MI)	41° 57.6 N	83° 15.4 W	
24	9063085	Toledo (US, OH)	41° 41.6 N	83° 28.3 W	
25	9063079	Marblehead (US, OH)	41° 32.6 N	82° 43.9 W	
26	9063063	Cleveland (US, OH)	41° 32.5 N	81° 38.1 W	
27	9063053	Fairport (US, OH)	41° 45.6 N	81° 16.9 W	
28	9063038	Erie (US, PA)	42° 09.2 N	80° 05.6 W	
29	9063028	Sturgeon Point (US, NY)	42° 41.5 N	79° 02.8 W	
30	9063020	Buffalo (US, NY)	42° 52.6 N	78° 53.4 W	

Table A6. List of permanent water level gauges located within the study region, and operated by theCanada's Department of Fisheries and Oceans (DFO) and the United States National Oceanic andAtmospheric Administration (NOAA).

Abbreviations: CA, Canada; US, United States; ON, Ontario; MI, Michigan; OH, Ohio; PA, Pennsylvania; NY, New-York.

System	Tributary Name	Water Quality Station ID
Segment		
SCR	St. Clair River Upstream	ECCC station at Point Edward (43.0048°N, -82.4155°W)
	St. Clair River Upstream	station at Port Huron (MIDEQ 740376; USEPA #3740376)
	Black River-MI	MIDEQ 740267; 21MICH_WQX-740267; USGS-04160075
	Belle-Pine Complex	UAL* from Clinton River
	St. Clair River Downstream	ECCC station at Port Lambton (42.6589°N, -82.5068°W)
	St. Clair River Downstream	Station at Algonac (MIDEQ 740016; USEPA #740016)
LSC	Clinton River	MIDEQ 500233
	Ruscom River	PWQMN 04001000302
	Thames River	PWQMN 04001308202 & WQMSD
	Sydenham River	PWQMN 04002701702 & WQMSD
DR	Detroit River Upstream	Station at Fort Wayne (21MICH-820414; USEPA #820017)
	Canard River	PWQMN 10000200202
	Turkey Creek	PWQMN 10000100302
	Rouge River	MIDEQ 820070
	Detroit River Downstream	GLWAP & 21MICH-820017
WB	Huron River-MI	OHEPA 580364
	Swan Creek	UAL* from Raisin River
	Raisin River	NCWQR
	Ottawa River	UAL* from Raisin River
	Maumee River	NCWQR
	Portage River	OHEPA 500510/NCWQR
CB	Sandusky River	NCWQR
	Huron River-OH	OHEPA 501030
	Vermilion River	NCWQR/OHEPA 501260/USGS 04199500
	Black River-OH	OHEPA 501510/USGS 04200500
	Rocky River	OHEPA 501790
	Cuyahoga River	NCWQR
	Chagrin River	OHEPA 502400
	Grand River-OH	NCWQR/502530
	Ashtabula-Conneaut Complex	USGS 04212945 (or UAL* from Grand River Ohio)
	Catfish Creek	PWQMN 16009700802/1600150104 & WQMSD
	Kettle Creek	PWQMN 16008701002/1600150103 & WQMSD
EB	Big Otter Creek	PWQMN 16010900802/1600150106 & WQMSD
	Erie-Chautaugua Complex	UAL* from 1. Cattaraugus Creek or 2. Grand River
	Cattaraugus Creek	NYDEC 01041302B/USGS 04213500
	Eighteenmile Complex	UAL* from 1. Cattaraugus Creek or 2. Grand River
	Big Creek	PWQMN 16012401102
	Lynn River	PWQMN 16015900302 & WQMSD
	Grand River-ON	PWQMN 16018403502 & WQMSD
	Nanticoke Creek	PWQMN 16016400102 & WQMSD
NR	Niagara River at Fort Erie	ECCC station at Fort Erie (station #ON02HA0045)

Table A7. Water quality stations used in deriving TP Loadings for 2003-2016. Abbreviations for thewater quality stations are explained in Table A8.

*UAL, the unit area load (UAL) extrapolation approach used in estimates of loads from unmonitored areas.

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Acronym	Description	Agency/Institution
ECCC	Upstream/Downstream Monitoring Program for rivers St. Clair and Niagara	Environment and Climate Change Canada
MIDEQ	Water Resources Division	Michigan Department of Environmental Quality
NCWQR	National Center for Water Quality Research	Heidelberg University
NYDEC	Division of Water	New York Department of Environmental Conservation
OHEPA	Division of Surface Water	Ohio Environmental Protection Agency
PWQMN	Provincial Water Quality Monitoring Network	Ontario Ministry of Environment, Conservation and Parks
STORET/NWIS	Water Quality Portal	United States Environmental Protection Agency (USEPA), USGS
USEPA	-	United States Environmental Protection Agency
USGS	Water Resources Division - NWIS	United States Geological Survey
WQMSD	Water Quality Monitoring and Surveillance Division	Environment and Climate Change Canada (ECCC)
WSC	Water Survey of Canada	Environment and Climate Change Canada (ECCC)

Table A8. Abbreviations for the water quality monitoring agencies listed in Table A7.

#	Location	Well ID	Latitude	Longitude	Ground	Well	Aquifer Type	Lithology of	# of TP
	(Lake Erie		(North)	(West)	Elevation	Depth		Aquifer	observations
	basin)				(m.a.s.l.)	(m.b.gr.)			
1		W0000099-1	42°1'53"	82°54'4"	192.58	39.32	Bedrock	Limestone	13
2		W0000358-1	42°2'16"	83°1'49"	177.85	19.20	Interface	Gravel, limestone	13
3	Western	W0000359-1	N/A	N/A	203.27	41.15	Bedrock	Limestone	13
4		W0000205-2	42°7'22"	82°36'10"	194.00	28.96	Overburden	Sand	4
5		W0000112-1	42°5'32"	82°51'16"	190.00	32.77	Interface	Silty till, Sand, Silt	12
6		W0000204-1	41°46'4"	82°40'14"	175.11	28.96	Bedrock	Limestone	13
7		W0000237-1	N/A	N/A	180.92	42.06	Bedrock	Shale	12
8		W0000236-1	N/A	N/A	189.81	50.9	Overburden	Gravel, Sand	1
9		W0000181-1	42°19'12"	81°56'11"	179.45	38.1	Interface	Gravel, Shale	11
10	Central	W0000249-1	42°25'10"	81°55'49"	200.09	35.66	Overburden	Gravel, Sand	13
11		W0000445-1	42°34'13"	81°31'41"	199.38	85.34	Bedrock	Shale	7
12		W0000185-1	N/A	N/A	212.05	42.98	Overburden	Sand	11
13		W0000452-1	42°42'31"	81°11'48"	210.58	60.96	Overburden	Sand, Gravel	11
14		W0000335-2	42°43'37"	80°53'2"	225.75	10.97	Overburden	Sand, Silty clay	3
15		W0000169-1	42°39'31"	80°30'38"	208.23	12.5	Overburden	Sand	3
16		W0000171-2	42°45'2"	80°25'12"	231.53	7.62	Overburden	Sand	3
17	Eastern	W0000271-1	42°50'4"	80°8'57"	205.12	41.15	Bedrock	Limestone	5
18		W0000178-1	42°56'40"	79°56'46"	216.365	25.91	Bedrock	Limestone	13
19		W0000289-1	42°54'4"	79°6'56"	186.34	4.5	Bedrock	Limestone	18

Table A9. Name, location, and characteristics of Canadian ground water monitoring wells* with reported concentrations of TP in the ground water and located within 8 km wide band of shoreline surrounding the lake boundaries.

* The Canadian monitoring well data was obtained from the Ontario Provincial Groundwater Monitoring Network: PGMN (PGMN, 2022); the sampling frequency is one sample per year.

Table A10. Mean over-lake precipitation (Y ; m ³ s ⁻¹), calculated volume-weighted mean total phosphoru	IS
(TP) concentration $(p^{PR}; mg L^{-1})$ in precipitation, and over-lake TP wet and total deposition rate	es
$(D^w, D^t;$ Metric Tonnes per Annum: MTA) in Lake St. Clair and Lake Erie for each water year	
during the study period $(2003 - 2016)$.	

Water	r Lake St. Clair				Lake Erie					
i cai	p^{PR}	Y	D^w	D^t	$p^{PR'}$	$p^{PR''}$	p^{PR} (mean)	Y	D^w	D^t
2003	0.0304	21.90	21.0	72.4	0.0280	0.0114	0.0197	580.3	360.5	1243.8
2004	0.0311	20.00	19.7	67.8	0.0152	0.0156	0.0154	564.3	274.9	948.4
2005	0.0321	22.40	22.7	78.3	0.0122	0.0090	0.0106	447.4	149.4	515.3
2006	0.0452	31.20	44.5	153.5	0.0090	0.0087	0.0089	722.3	201.9	696.5
2007	0.0267	26.40	22.2	76.7	0.0287	0.0136	0.0211	799.9	532.6	1837.6
2008	0.0066	23.00	4.8	16.6	0.0051	0.0175	0.0113	658.6	235.5	812.6
2009	0.0216	25.30	17.2	59.4	0.0317	0.0156	0.0237	577.1	430.4	1484.9
2010	0.0091	32.60	9.3	32.2	0.0062	0.0122	0.0092	825.4	239.2	825.3
2011	0.0067	27.30	5.8	19.9	0.0127	0.0167	0.0147	604.0	279.2	963.2
2012	0.0087	29.80	8.2	28.3	0.0080	0.0110	0.0095	684.3	205.1	707.5
2013	0.0182	26.30	15.1	52.1	0.0120	0.0202	0.0161	580.3	295.0	1017.6
2014	0.0098	29.10	9.0	30.9	0.0050	0.0128	0.0089	702.0	198.1	683.3
2015	0.0064	28.80	5.8	20.0	0.0068	0.0153	0.0110	656.2	228.6	788.7
2016	0.0067	30.00	6.3	21.9	0.0115	0.0238	0.0177	708.2	395.5	1364.5
Average:	0.0185	26.7	15.1	52.1	0.0137	0.0145	0.0141	650.7	287.6	992.1

Note: Two wet precipitation stations for Lake Erie, at Point Pelee and Rock Point (Fig. 1; Appendix A5), are indicated with symbols <'> and <''>, respectively.

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Figure A1. Map of Lake Erie sub-watersheds (modified from the map courtesy to the Ohio Department of Natural Resource - ODNR map library). The eastern most sub-watershed with identification code 04120004, named "Niagara" and discharging directly to the Niagara River, was not included into the water and phosphorus mass budgets.

Appendix A1. Defining an appropriate temporal scale for a mass budget construction.

To construct a mass budget based on input-output approach for a given conservative substance (e.g. water) in any system with a significant volume (V) compared to its flow-through rate (Q) and resulting a high water storage capacity or mean water residence time (R_w ; $R_w = V/Q$), it is important to construct a budget over the period of time appropriate for the transport time scales (e.g. larger than water age for a water parcel at any location in the lake) and which is long enough to minimize the effects of initial conditions, and keep the effect of initial conditions on the results as small as possible.

 R_w , a water storage capacity, or also known as a mean water residence time, in Lake Erie System (LES) is primarily determined by Lake Erie's R_w (Bolsenga and Herdendorf, 1993: 1000 days or ~2.8 years; this study: 944 days or ~2.6 years, Table 4) because of negligible value of R_w for Lake St. Clair (~ 9 days; Bocaniov et al., 2019), the St. Clair River (0.88 day; Griffiths et al., 1991) and Detroit River (0.83 day; Derecki, 1984). Assuming Lake Erie as an ideally mixed continuous steering tank reactor (CSTR) and considering a Lake Erie's $R_w = 2.8$ years, then 4.6 times of τ or 13 years would be needed for the outflow to contain a 1% of the initial lake concentration. Therefore, in this study we decided to use a period of 14 years which is larger than 4.6 times R_w , so that the lake concentrations will be in full equilibrium (>99%) with the external loads. Over 14 years when the steady state is reached, the in-lake P concentrations must be in equilibrium with the loads, so the left sides in Equations 8 and 10-12 (Table 1) will be zero, so the internal net P load can be estimated for each specific segment.

Appendix A2. *The run-off from the Thames, Sydenham, and Grand Ontario rivers.*

The run-off from the three largest sub-watersheds (the Thames, Sydenham, and Grand Ontario rivers) was estimated using the following approach. For these sub-watersheds, the most downstream gauging stations were relatively far away from the river outlets with significant changes occurring in elevations between river outlet and monitored areas (up to 360 m; for example, the Grand Ontario River: elevations at the source and at the mouth are 525 and 174 m above sea level, respectively) which may affect along-river precipitation pattern. For these three sub-watersheds we calculated the total runoff using the exponential relationship (Eq. A2.1):

$$R = R_g \cdot K^S \tag{A2.1}$$

$$K = \frac{A}{A_g} \tag{A2.2}$$

$$S = \frac{\log_{10}\left(\frac{R_1}{R_2}\right)}{\log_{10}\left(\frac{A_1}{A_2}\right)} \tag{A2.3}$$

where, *R* is the total runoff into the lake, R_g is the runoff at the furthermost downstream gaging station in the basin, *K* is the ratio of the total area of the drainage basin (*A*) to the drainage area above the gaging station (A_g), *S* is the basin exponent, R_1 and R_2 are the gauged runoffs from downstream and upstream gauges, A_1 and A_2 are the gauged areas corresponding to R_1 and R_2 .

Appendix A3. Approach to close the water budget.

The constructed budget revealed the negative residuals of 6 and 52 m³ s⁻¹ for lakes St. Clair and Erie, respectively. Though, this is a small amount accounting for less than 1% of the combined flow for connecting channels and tributary, it must be balanced. It was done by correcting flows with the largest uncertainties in estimates and using them as a target for budget closure. Lake water levels as well as flows in connecting channels and water diversions are relatively accurately monitored using a network of land-and shoreline based in-situ gauge measurements. Consumptive water use is also strictly regulated and monitored (GLC, 2019). Precipitation is also measured via a network of land-based and shoreline-based rain gauges. The reported direct groundwater water discharge into the lake is small accounting for not more than 12 m³ s⁻¹. The most uncertain element is evaporation which is a simulated value and not a measured one. Moreover, in a recent study, Shao et al. (2020) measured evaporation rates at two stations in the western basin from Sept. 2011 to May 2016 and reported lower values for mean annual evaporation compared to those derived from the GLERL hydrologic database (GLERL, 2020a) which could suggest that the evaporation rates in the database may slightly overestimate the open-lake evaporation results. Therefore, to close the water budget we adjusted our mean annual evaporation rates from 37 to 31 m³ s⁻¹ and from 763 to 715 m³ s⁻¹ in lakes St. Clair and Erie, respectively (Table 6; Fig. 2b).

Appendix A4. Water budget.

Over-lake precipitation data were obtained from the NOAA Great Lakes Environmental Research Laboratory (GLERL) Great Lakes Monthly Hydrologic Data base (GLERL, 2020a). Over-lake evaporation data were also obtained from GLERL (2020a), where it is calculated using daily simulations from NOAA-GLERL's one-dimensional Net Basin Supply model (Hunter et al., 2015).

Daily flow rates for the major tributaries were downloaded from the USGS National Water Information System (USGS, 2019) for sites in the US and from the National Water Data Archive HYDAT (HYDAT, 2020) for sites in Canada (Tables A4 & A5), scaled to the entire sub-watershed area to include the unmonitored areas (Tables A2 & A3) based on area-weighted estimates for the closest monitored areas in most cases. The approach for the three largest sub-watersheds in Canada (the Thames, Sydenham, and Grand Ontario rivers) is described in Appendix A2.

For the connecting channels, the St. Clair and Detroit rivers, we used the USGS station 04159130 at Port Huron and 04165710 at Fort Wayne, respectively (Table A5: #1 and #6) for October 2008 to 2016 and onwards. For October 1 2002 to September 2008, we calculated daily flows based on daily mean water levels and stage-fall-discharge relationships (Fay and Kerslake, 2009) similar to Scavia et al. (2019a, b). To estimate the daily flows for the outflow, the Niagara River (Fig. 1), we used USGS station 04216000 Niagara River at Buffalo NY; Table A5: #36).

Our estimate of direct groundwater inflow to Lake Erie was based on the results of the fullyintegrated surface water-groundwater model (Xu et al., 2021) that estimated an annual average rate to be 11.9 m³ s⁻¹.

The outflow through the Welland Canal was estimated as the reported data by WSC (Water Survey of Canada; WLF, 2020) for station 02HA019 (Table A4: #30). Apart from the Welland Canal diversion, there are two small diversions bringing Lake Huron water to the system, bypassing the St. Clair River: water intakes for Detroit, Michigan and London, Ontario. London diverts about 3 m³ s⁻¹ from Lake Huron and returns it to Lake St. Clair via the Thames River (Quinn and Edstrom, 2000) as treated sewage.

The sewage outfall is upstream of the water flow gauging station 02GE003 (Table A4: #9) which means that this diverted amount of water is already accounted for by this gauging station. Detroit also withdraws water from Lake Huron via its Lake Huron WTP water intake and returns it to Lake St. Clair. The rated capacity of the Lake Huron WTP is 17.5 m³ s⁻¹ (DWSD, 2019). However, due to current conditions of treatment facilities and water demand, the maximum and average daily flows of this WTP is only 12.3 m³ s⁻¹ and 7.0 m³ s⁻¹, respectively (2003 to 2015 data: DWSD, 2019).

Consumptive water use was defined as the difference between the amount of water withdrawn for all purposes including but not limited to municipal supply, agricultural and industrial uses and the amount of water returned to the system on the annual base, and was estimated for Lake Erie system based on the annual report of the Great Lakes Regional Water Use Database representing water use data 2003 through 2016 and prepared by the Great Lakes Commission (GLC, 2019)

Annual net changes in storage of lakes St. Clair and Erie were calculated as the product of the measured changes in lake stage and surface area of each lake. In large shallow lakes, as our lakes St. Clair and Erie, the water levels are highly sensitive to the wind set-ups (storm surges) and surface seiches which can cause great variability in lake stage at different parts of a lake, we used available measurements of mean daily water levels at multiple locations around lakes obtained from a permanent network of water level gauges operated by the Canada's Department of Fisheries and Oceans (DFO) and the United States National Oceanic and Atmospheric Administration (NOAA) and listed in Table A6.

Appendix A5. Estimates of atmospheric total phosphorus (TP) deposition.

Within the Great Lakes Basin, Environment and Climate Change Canada (ECCC) maintains a nine-station network of atmospheric deposition stations, named the Great Lakes Precipitation Network (GLPN). Within GLPN, Canada operates one station located on Lake St. Clair, and two stations located on Lake Erie. The St. Clair site is located on the east shore of Lake St. Clair (Fig. 1). The two stations on Lake Erie are located at Point Pelee and Rock Point. Point Pelee station is located at the southwestern end of Lake Erie near the Point Pelee, while Rock Point station is located on the northeastern shore of Lake Erie (Fig. 1). Station locations and sampling protocols are described in Chan et al. (2003). All three stations are downwind of the major industrial centers and heavily industrialized regions of the Detroit, Windsor, Sarnia and Cleveland areas. The sampling sites are located in the close proximity of lake shores and deemed to be regionally representative of each lake atmospheric conditions. These three stations measure total phosphorus concentration in wet/frozen precipitation and can be used for estimates of wet deposition.

The Volume Weighted Mean (p_i^{PR}) concentration in precipitation for each water year for segment *i* (where *i* can stand for either Lake St. Clair or Lake Erie) was calculated as the sum of the masses of TP measured in each sample divided by the sum of the volumes of all the samples of precipitation collected for all samples during the given year:

$$p_i^{PR} = \frac{\sum_{j=1}^{j=N} C_j \cdot V_j}{\sum_{j=1}^{j=N} V_j}$$
(A5.1)

where, C_i is the concentrations in the sample (mg L⁻¹), V_j is the volume of the sample *j* (L), *j* is a rain sample, and *N* is the total number of rain samples taken within a given year.

To determine the wet deposition flux (F_i^w ; mg m⁻² year⁻¹) for segment *i*, we used the annual mass of TP collected at each site, divided by the area (A; m²) of the collection bucket (0.0314 m²), corrected for the rain water collection area and difference between the actual precipitation and those captured by the rain collectors. Supporting Information "A" for Bocaniov et al. (2023): Long-term phosphorus mass-balance of Lake Erie (Canada-USA) reveals a major contribution of in-lake phosphorus loading. *Ecological Informatics (print ISSN: 1574-9541; online ISSN: 1878-0512)*.

$$F_{i}^{W} = \frac{p_{i}^{PR} \cdot \sum_{j=1}^{j=N} V_{j}}{A}$$
(A5.2)

For Lake Erie, the average F_{LE}^{w} was calculated as the mean of two site-specific values measured at Pelee and Rock Point sampling stations ($F_{LE}^{w'}$ and $F_{LE}^{w''}$, respectively). For Lake St. Clair, F_{LSC}^{w} was taken as for the station St. Clair.

Because of one or two contaminated or missing samples within each year, the wet deposition flux was corrected for the difference between collected precipitation and actual over-lake precipitation:

$$F_i^{w,cor} = \frac{Y_i \cdot A}{\sum_{j=1}^{j=N} V_j}$$
(A5.3)

The annual wet deposition rate $(D^w; MTA)$ over the entire lake was estimated as:

$$D_i^w = F_i^{w,cor} \cdot A_i \cdot 10^{-9}$$
(A5.4)

where, *A* is a lake area (m²; see Table 4: Lake St. Clair: $A = 1114 \times 10^6 \text{ m}^2$; Lake Erie: $A = 25657 \times 10^6 \text{ m}^2$), 10^{-9} is a coefficient to convert milligrams (mg) into metric tonnes (MT).

Based on the measurements of wet and dry TP deposition in the agricultural regions within the continental United States (Anderson and Downing, 2006), the dry deposition rates over the entire lake $(D^d; MTA)$ were assumed to be 2.45 times of the D^w rates. The total deposition $(D^t; MTA)$ determined as the sum of D^w and D^d , was calculated by multiplication of D^w by 3.45.

Appendix A6. On estimates of the TP sedimentation rates in Lake Erie.

The current appendix provides a comparison of the sedimentation rates in Lake Erie estimated using the apparent settling velocities reported by Chapra et al. (2016) with those reported by literature.

Our sedimentation rates of TP in the western (3188 MTA) and central (5032 MTA) basins are reasonable and consistent with the previous studies. For example, Kemp et al. (1977) estimated the sediment loads to the western and central basins as 2.681 and 5.866 million Metric Tonnes per Annumn (MTA) per year, respectively. The field observations (Williams et al., 1976) suggest the average values of 1,150 and 0.900 μ g of P per g dry weight of sediments in the depositional areas of the western and central basins, respectively. This will give us annual TP sedimentation rates of 3,100 and 5,300 MTA in the western and central basins, respectively.

Our TP sedimentation rate for the eastern basin maybe smaller as can be predicted from the sediment budget for this basin with the sedimentation rate of 5.8 million MTA (Kemp et al., 1977; Mortimer, 1987). Such a sedimentation rate assumes the deficit of 4 million MT (Mortimer, 1987) and sediment budget that is not closed. If someone can assume that the sedimentation rates were somehow overestimated by the amount of the deficit (4 million MTA) so it is not 5.8 million but 1.8 million MTA, then all inputs are balanced with the outputs within the eastern basin. This number (1.8 million MT) is closer to our estimate of the TP sedimentation of 1500 MTA. According to Williams et al. (1976) the average value of 1000 μ g of P per g dry weight of sediments in the depositional areas of the eastern basin. This would mean a TP sedimentation flux of 1800 MTA. The difference between this and our estimate can be attributed to the arrival and establishment of mussels and their ability to filter the particulate matter and redirect fluxes of phosphorus from the particulate pool to the dissolved pool.

Our estimate TP sedimentation areal rate in the eastern basin is 0.233 gP m⁻² yr⁻¹ (Table 8). Bloesch (1982) reported mean summer (Jun 28 to Sept 8, 1978) deposition rates of Particulate Phosphorus (PP) for his offshore station (depth = 40 m; northern side of the basin) at two depths, 3 m above bottom (hypolimnetic fluxes) and just below epilimnion (epilimnetic fluxes), as 5.44 mgP m⁻² d⁻¹ and 1.3 mgP m⁻² d⁻¹, respectively. He observed that epilimnetic fluxes of PP were 4-20 times less than the PP fluxes measured in the hypolimnion due to effect of frequent resuspension events at the bottom of the lake. This means that the measured sedimentation rates near the bottom of the eastern basin can be significantly overestimated. This may suggest that the high sedimentation rates of 5.8 million MTA finegraned sediments observed by Kemp et al. (1977) in the eastern basin were overestimated by a factor of 3 as they were not corrected for the bottom resuspension effects, and in this case no deficit of 4 million MT of fine-grained sediments exists in the eastern basin. Moreover, if Bloesch' (1982) summer averaged epilimnetic sedimentation flux of PP (1.3 mgP m⁻² d⁻¹ = 0.0013 gP m⁻² d⁻¹) is assumed to be the same for the entire year (0.0013 gP m⁻² d⁻¹ x 365.25 d yr⁻¹ = 0.475 gP m⁻² d⁻¹) then it will equal to 0.475 gP m⁻² d⁻¹ that is only 2 times higher than our annual rate of 0.233 gP m⁻² d⁻¹ (Table 8), and it is reasonable to assume that the mean summer sedimentation rates are higher than the mean annual rates, and that the arrival of mussels (~1990) resulted in clearing of the water column and reduced standing crop of the detrital carbon and phytoplankton biomass not only nearshore but also offshore (e.g. see Bocaniov et al., 2014, for the nearshore offshore differences in the phytoplankton biomass, and for the effect of the advection of the nearshore waters with lower Chl-a on the offshore biomass of Chl-a). So, the pre-mussel (e.g. 1978 rates) sedimentation rates should be smaller than the after-mussel sedimentation rates (e.g. 2003 to 2016).

There are other indirect indications that the offshore sedimentation rates of TP are relatively small. For example, the biomass of mussels in the deep offshore areas of the eastern basin is large but the population is composed from the large-sized older individuals with the absence of smaller-size

Appendix A7. Total phosphorus (TP) mass balance.

Tributaries and Connecting Channels – All tributary and connecting channel nutrient loads were estimated using the Water Regressions on Time, Discharge, and Season (WRTDS) method (Hirsch et al., 2010) based on observed concentrations and flow data. To estimates P loads with three connecting channel flows (rivers St. Clair, Detroit and Niagara) where TP concentrations are not correlated with channel discharges and WRTDS method cannot be applied, similar to other previous studies (Scavia et al., 2019a, b), we used the Generalized Additive Model (Wood, 2011; Wood, 2017) to approximate daily TP concentrations as a function of time, with the smoothing parameters selected by the Restricted Maximum Likelihood (REML) method. The estimated daily TP concentration and associated standard error were multiplied by daily discharge to estimate daily loading. The daily estimates were then aggregated to estimate annual loadings.

TP concentration data were obtained from the United States Water Quality Portal (WQP, 2019), the Provincial (Stream) Water Quality Monitoring Network (PWQMN, 2019), Michigan Department of Environmental Quality (MDEQ), the Great Lakes Intake Program (GLIP, 2019) managed by the Ontario Ministry of the Environment, Conservation and Parks (MECP), and Environment and Climate Change Canada (ECCC; communication with D. Burniston and A. Dove, 2018).

For the St. Clair River, TP loads were estimated at the river source (Point Edward, Fig. 1: site #1; Port Huron, Fig. 1: site #2), and two sites close to the river mouth (Port Lambton, Fig. 1: site #3; Algonac, Fig. 1: site #4). At some stations (e.g. St. Clair River at Port Huron), similar to Scavia et al. (2019a, b), we removed extreme outliers defined here as concentrations higher than the mean plus four times the standard deviation (SD) and representing less than 0.02% of data.

Due to lower resolution of the water quality and nutrient monitoring activities at the river source (Point Edwards; Port Huron) compared to the timing and duration of the short-lived episodic events (shoreline erosion and resuspension of sediments in the nearshore due to storm surges, searches and surface waves; see Scavia et al., 2019a, b), some of the TP load is missed. To estimate it we calculated as Supporting Information "A" for Bocaniov et al. (2023): Long-term phosphorus mass-balance of Lake Erie (Canada-USA) reveals a major contribution of in-lake phosphorus loading. *Ecological Informatics (print ISSN:1574-9541; online ISSN: 1878-0512)*.

the difference of the TP load at the St. Clair River downstream location (average of loads calculated at both sites, Point Edwards and Port Huron) and a sum of estimated loads for the St. Clair River upstream location and for the direct river sub-watersheds. The load of the St. Clair River to Lake St. Clair was calculated as the average of loads calculated for both sides of the river, at Port Lambton (Ontario; Fig. 1: site #3) and Algonac (Michigan; Fig. 1: site #4).

For the LSC TP load to Detroit River we used station #21MICH-820414 (USEPA #820017) at the lake outflow (Table A7; Fig. 1: site #5). We estimated the total annual load from the Detroit River to the western basin of Lake Erie as a sum of this LSC load and loads from Detroit River sub-watersheds. This is because water quality measures at the downstream monitoring station (#21MICH-820017; Table A7) can be biased due towards frequent episodes of Lake Erie surface seiches and wind-driven water setups transmitting lake water upstream from the lake to the river.

Point Sources - Point sources identified here are only those with direct inputs downstream of the river mouth monitoring stations. Those discharging upstream of the monitoring stations were included in the tributary loads. Direct annual point TP loads from 2003 to 2016 for each sub-watershed were from Scavia et al. (2019a, b) for the St. Clair River, Lake St. Clair, and Detroit River. For the three basins of Lake Erie, direct loads were those in Maccoux et al. (2016). The latter study reported those loads for only 2003 to 2013. Because there was very little inter-annual variability, we assumed the loads for 2014 to 2016 to be the same as averaged 2003-2013 values. This assumption should not affect results or conclusions because direct TP loads are a very small compared to the total watershed loads (Maccoux et al., 2016; Scavia et al., 2019a, b).

Groundwater - The volume of direct groundwater to Lake Erie is small (11.9 m³ s⁻¹; Xu et al., 2021) accounting for less than 1% of the combined over-lake precipitation and basin runoff or less than 0.25% of the Detroit River inflow to the lake. In terms of phosphorus concentrations, the direct ground water discharge into Lake Erie can be characterized by chemical composition of ground water in monitoring wells located within 8 km wide band surrounding the lake, because this zone, along with the lake bed, was suggested as the zone most important for ground water – lake water exchange processes

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(Xu et al., 2021). The Canadian monitoring well data was obtained from the Ontario Provincial Groundwater Monitoring Network (PGMN) Program (PGMN, 2022) and the U.S monitoring well data was obtained from the USGS National Ground-Water Monitoring Network (USGS, 2022). Within such an 8 km wide band zone, on the Canadian portion of the lake, we identified 19 ground water monitoring wells reporting total phosphorus concentrations (Table A9), and there were no wells found on the USA portion of the lake with reported measurements of water chemistry. Therefore, our estimates of TP concentrations in the direct inputs of deep groundwater were based on Canadian sources. Because of the low frequency of data (one sample per year) and to extend the number of observations, we extended the period of observations to all available data, from 2002 to 2019. To reduce the importance of a few large outliers, present in the dataset, that might result from local contamination, we decided to use a pooled dataset for the entire basin and base our analysis on the most typical value that is median rather than using an average value.

Over-lake atmospheric deposition - Over-lake atmospheric wet deposition of TP for each water year were based on the product of volume-weighted rain water mean concentration (p_i^{PR}) and precipitation. p_i^{PR} was calculated based on the TP concentrations in precipitation by Environment and Climate Change Canada (ECCC) and over-lake precipitation amount obtained from GLERL (2020a) (see Appendix A5 for more details). Total over-lake deposition was assumed to be 3.45 times wet deposition based on estimates of wet and dry depositions for the continental United States (Anderson and Downing, 2006; see Appendix A5 for more details).

Outputs via the Niagara River and Welland Canal - TP losses via the Niagara River was estimated with GAM based on the data for river outflow and nutrient concentrations collected at the ECCC upstream monitoring site at Fort Erie (#ON02HA0045; Table A7; Fig. 1: site #6). Due to close proximity of the sources of the Welland Canal and the Niagara River to each other (Fig. 1), it was assumed that the TP concentrations in the lake water outflowing through the water diversion and lake outflow are similar. Therefore, the TP loss via the water diversion was estimated as the proportion of the Niagara River TP load related to the differences in flows.

Consumptive use - Contribution of the water consumptive use to the loss of TP was calculated as the annual rate of water taken from the lake and not returned back multiplied by the average annual segment-specific P concentration.

Sedimentation - A certain fraction of a lake's TP loading is retained permanently via settling and permanent burial. While net sedimentation is a delicate balance between sedimentation and resuspension rates, over an annual cycle, it has been parametrized for each of the three Lake Erie basins including Lake St. Clair using long-term, mass balance dynamic modeling (e.g. Chapra et al., 2016). To estimate the P removal in each basin due to net sedimentation, we used the mean annual basin-specific concentrations of TP and apparent settling velocities normalized to the basin surface area for post-1990 rates as in Chapra et al. (2016; see also Tables 1, B5, B7, B9 to B11).

Inter-basin nutrient exchange (bulk mixing) – To characterize the inter-basin nutrient fluxes due to diffusion we used the bulk diffusion coefficients parameterized in Chapra et al. (2016; see their Table 3) and mean annual basin-specific concentrations in the adjacent basins (Tables 1, B9 to B11).

Change in storage - The TP losses because of changes in lakes St. Clair and Erie volumes were calculated as the product of the mean annual difference in lake volume and its average lake-wide annual concentration. For Lake Erie, we used the mean of the observed median values for all three basins from the two spring and summer lake-wide monitoring cruises representing both the isothermal (April) and stratified (August) conditions (GLNPO, 2020). The median values were chosen as a good representation of the most typical nutrient conditions as they are not inflated with potential sources of contamination, with a similar approach undertaken by Chapra et al. (2016). To estimate the mean TP concentration for Lake St. Clair over the study period we used observations of TP concentrations at the lake outflow (#21MICH-820414, USEPA #820017, Fig. 1: site #5, Table A7). Preliminary comparisons of this concentration with the lake averaged concentration determined from all lake loads and water flows, returned similar results, validating the estimated value.

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Internal P load – It was determined as the difference between means of all annual loss and input

fluxes of TP averaged over 2003 to 2016 using Equation 8 (Table 1) for Lake St. Clair and Equations 10-12 (Table 1) for the three basins of Lake Erie.

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